Introduction

The Oregon Department of Geology and Mineral Industries (DOGAMI) has been identifying and mapping the tsunami inundation hazard along the Oregon coast since 1994. In Oregon, DOGAMI manages the National Tsunami Hazard Mitigation Program, which has been administered by the National Oceanic and Atmospheric Administration (NOAA) since 1995. DOGAMI's work is designed to help cities, counties, and other sites in coastal areas reduce the potential for disastrous tsunami-related consequences by understanding and mitigating this geologic hazard. Using federal funding awarded by NOAA, DOGAMI has developed a new generation of tsunami inundation maps to help residents and visitors along the entire Oregon coast prepare for the next Cascadia Subduction Zone (CSZ) earthquake and tsunami as well as for far-travelled, or "distant" tsunamis.

The "Ring of Fire," also called the Circum-Pacific belt, is the zone of earthquake activity surrounding the Pacific Ocean. It is an arc stretching from New Zealand, along the eastern edge of Asia, north across the Aleutian Islands of Alaska, and south along the coasts of North and South America (Figure 1). The Ring of Fire is located at the borders of the Pacific Plate and other major tectonic plates. The Pacific Plate is colliding with and sliding underneath other plates creating subduction zones that eventually release energy in the form of an earthquake rupture. This rupture causes a vertical displacement of water that creates a tsunami. When these events occur around the Ring of Fire and but not directly off the Oregon coast, they take more time to travel the Pacific Ocean and arrive onshore (Figure 2). Distant earthquake/tsunami events have occurred elsewhere on the planet: for example, offshore Chile in 1960 and 2010, offshore Alaska in 1964, near Sumatra in 2004, and offshore Japan

Historically, about 28 distant tsunamis have been documented by Oregon tide gages since 1854. The most severe was generated by the 1964 M9.2 Prince William Sound earthquake in Alaska. Oregon was hit hard by the tsunami, which killed four people and caused an estimated 750,000 to 1 million dollars in damage to bridges, houses, cars, boats, and sea walls. The greatest tsunami damage in Oregon did not occur along the ocean front as one might expect, but in the estuary channels located farther inland. Of the communities affected, Seaside was inundated by a 10 foot tsunami wave and was the hardest hit. Tsunami wave heights reached 10 to 11.5 feet in the Nehalem River, 10 to 11.5 feet at Depoe Bay, 11.5 feet at Newport, 10 to 11 feet at Florence, 11 feet at Reedsport, 11 feet at Brookings, and 14 feet at Coos Bay

Prince William Sound event, and the second scenario represents a inundation extent for the two scenarios at select locations on this map. hypothetical maximum event. This maximum event is the same model used by the U.S. Geological Survey (USGS) in their 2006 tsunami hazard assessment of Seaside (TPSW, 2006). This model uses extreme fault model parameters that result in maximum seafloor uplift; nearly twice as large as in the 1964

earthquake. The selected source location on the Aleutian chain of islands also shows higher energy directed toward the Oregon coast than other Alaskan source locations. For these reasons the hypothetical "Alaska Maximum" scenario is selected as the worst case distant tsunami scenario for Oregon. For more detailed information on the tsunami models and methodologies used, please see DOGAMI Publication Special Paper 43 (Witter and others, 2011).

This tsunami inundation map displays the output of computer models representing the two selected tsunami scenarios; Alaska M9.2 (1964) and the Alaska Maximum. All tsunami simulations were run assuming that prevailing tide was static (no flow) and equal to Mean Higher High Water (MHHW) tide; MHHW is defined as the average height of the higher high tides observed over an 18-year period at the Port Orford tide gauge. The map legend depicts the respective amounts of deformation and the earthquake magnitude for these two scenarios. Figure 3 shows the cumulative number of buildings inundated

The computer simulation model output is provided to DOGAMI as millions of points with values that indicate whether the location of each point is wet or dry. These points are converted to wet and dry contour lines that form the extent of inundation. The transition area between the wet and dry contour lines is termed the Wet/Dry Zone, which equates to the amount of error in the model when determining the maximum inundation for the each scenario. Only the Alaska Maximum Wet/Dry Zone is shown on this map.

This map also shows the regulatory tsunami inundation line (Oregon Revised Statutes 455.446 and 455.447), commonly known as the Senate Bill 379 line. Senate Bill 379 (1995) instructed DOGAMI to establish the area of expected tsunami inundation based on scientific evidence and tsunami modeling in order to prohibit the construction of new essential and special occupancy structures in this tsunami inundation zone (Priest, 1995).

Time Series Graphs and Wave Elevation Profiles: In addition to the tsunami scenarios, the computer model produces time series data for "gauge" locations in the area. These points are simulated gauge stations that record the time, in seconds, of the tsunami wave arrival and the wave height observed for each time interval. It is especially noteworthy that the greatest wave height and velocity observed are not necessarily associated with the first tsunami wave to arrive onshore. Therefore evacuees should not assume that the tsunami Alaska-Aleutian Model Specifications: DOGAMI modeled two distant event is over until the proper authorities have sounded the all-clear signal at earthquake and tsunami scenarios involving M9.2 earthquakes originating the end of the evacuation. Figure 4 depicts time series data for the map plate near the Gulf of Alaska. The first scenario attempts to replicate the 1964 area. Figure 5 (profiles A-A' and B-B') depicts the overall wave height and

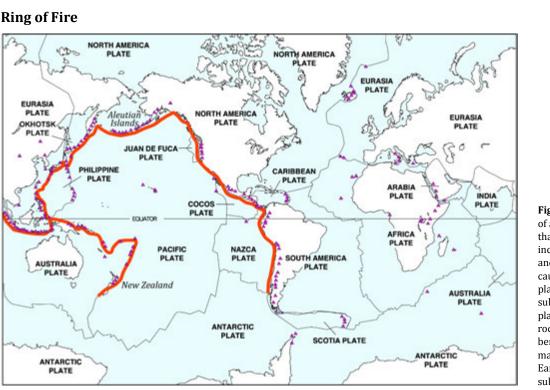


Figure 1: The "Ring of Fire" is a zone of active earthquakes and volcanoes that rings much of the Pacific Ocean, including the Oregon coast. Volcanoes and earthquakes on this ring are caused by the movements of tectonic plates. One type of movement is called subduction — when thin, oceanic plates such as those that compose the rock beneath the Pacific Ocean sink beneath thicker, lighter plates that make up continental plates. Earthquakes that occur as a result of

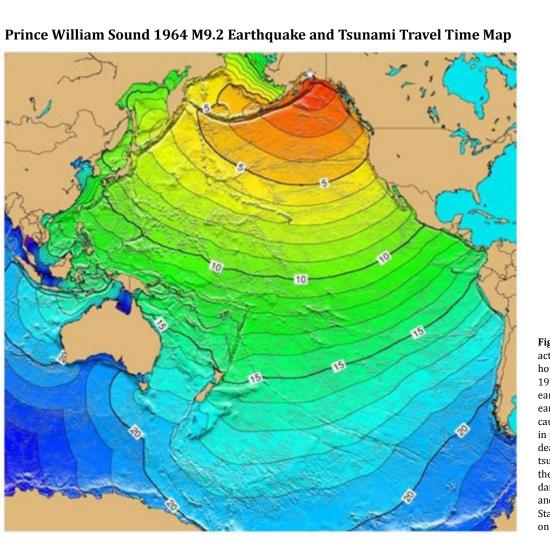
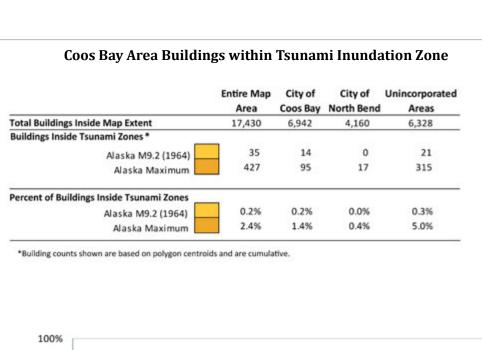
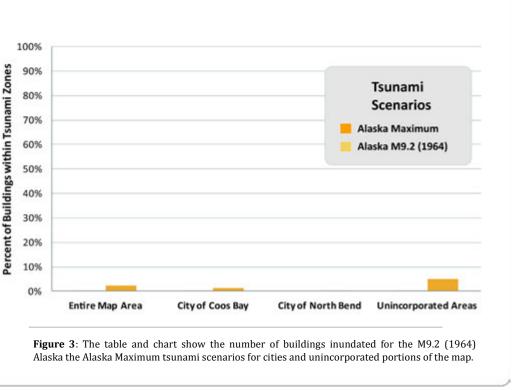
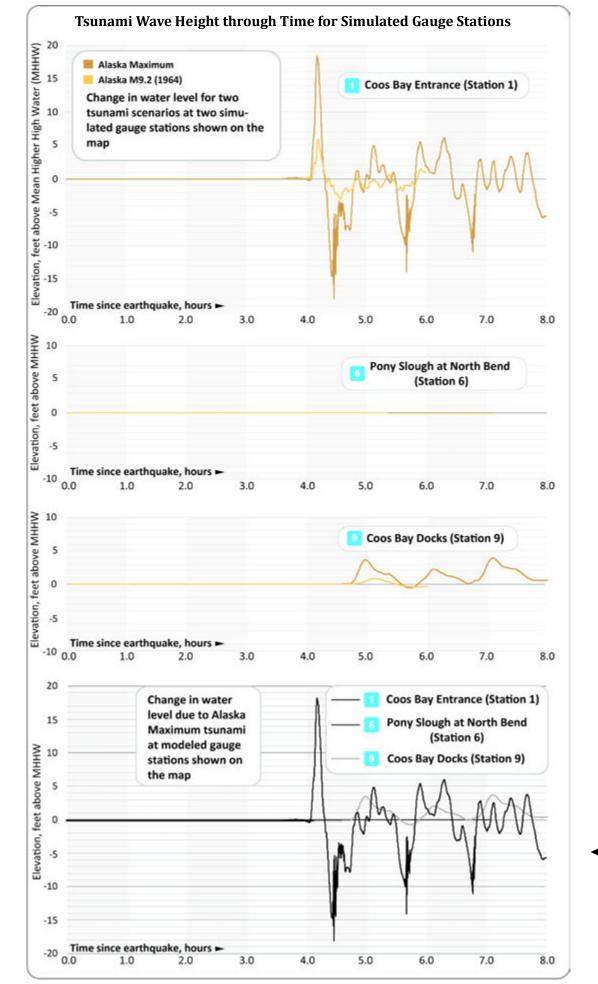
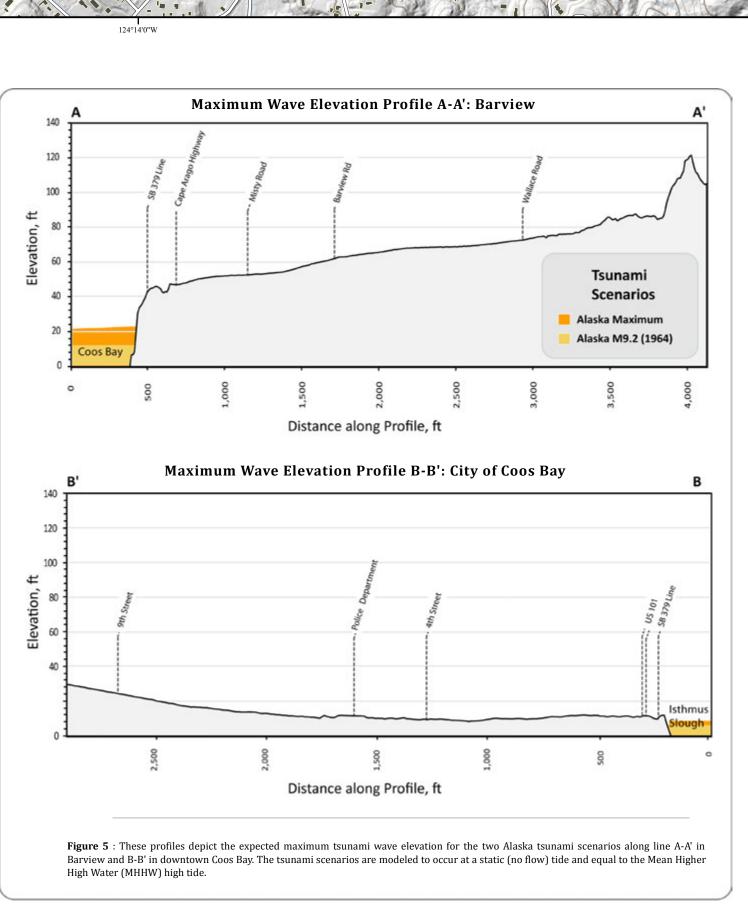


Figure 2: This image depicts the hours, around Pacific Rim from the 1964 Prince William Sound earthquake. This magnitude 9.2 caused 125 deaths and \$311 million in property loss, \$84 million and 106 deaths in Alaska (NGDC/WDC). The tsunami devastated many towns along the Gulf of Alaska. Left serious damage in British Columbia, Hawaii, and along the west coast of the United States left 15 dead, and was recorded on tide gages in Cuba and Puerto Rico.

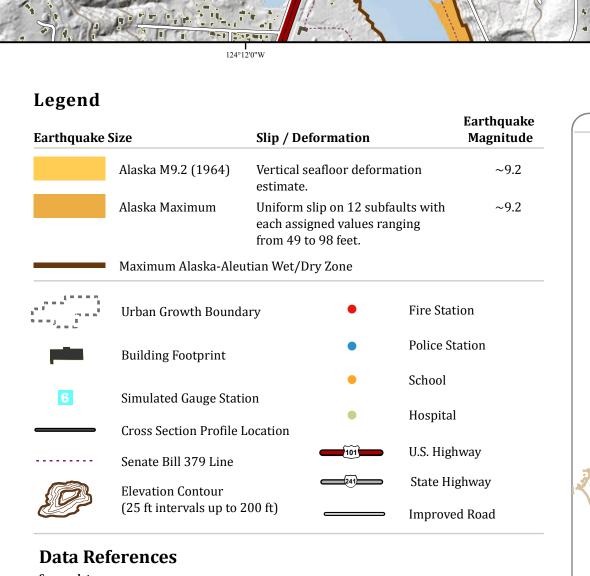






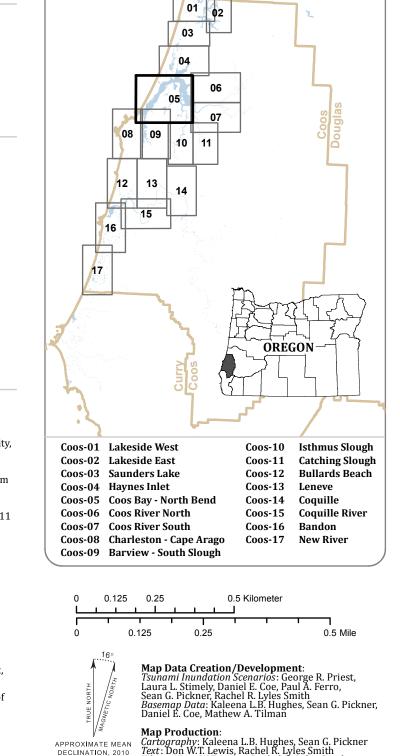


◀ Tsunami Wave Height over Time for Simulated Gauge Stations Figure 4: Top three charts depict the tsunami waves as they arrive at three selected reference points (simulated gauge stations). They show the change in wave heights for the two Alaska tsunami scenarios over an 8-hour period. The model predicts the first tsunami wave will arrive at the entrance to Coos Bay in approximately 4 hours; waves will continue to occur for up to 8 hours after the earthquake. Wave heights vary through time, and the first wave will not necessarily be the largest as waves interfere and reflect off local topography and bathymetry. Bottom chart depicts the change in wave height for the Alaska Maximum tsunami scenario only. Modeled wave heights, arrival times, and wave durations can help emergency response personnel plan for a tsunami.



Urban growth boundaries (2010) were provided by the Oregon Department of Land Conservation and Development (DLCD). Transportation data (2008) were provided by Coos County. Lidar data are from DOGAMI Lidar Data Quadrangles LDQ-2009-42124G4-Sixes and LDQ-2009-South Coast **Coordinate System**: Oregon Statewide Lambert Conformal Conic, Unit: International Feet, Datum: North American Datum 1983 HARN. Graticule shown with geographic coordinates (latitude/longitude). References:
National Geophysical Data Center / World Data Center (NGDC/WDC) Global Historical Tsunami Database,
Boulder, CO, USA. [http://www.ngdc.noaa.gov/hazard/tsu_db.shtml]. Priest, G. R., 1995, Explanation of mapping methods and use of the tsunami hazard maps of the Oregon coast, Oregon Department of Geology and Minerals Industries Open-File Report O-95-67, 95 p. Tsunami Pilot Study Working Group (TPSW), 2006, Seaside, Oregon tsunami pilot study — modern-ization of FEMA flood hazard maps: U.S. Geologi¬cal Survey Open-File Report 2006-1234, 90 p. + 7 app. [http://pubs.usgs.gov/of/2006/1234/]. Witter, R.C., Zhang, Y., Wang, K., Priest, G.R., Goldfinger, C., Stimely, L.L., English, J.T., and Ferro, P.A., 2011, Simulating tsunami inundation at Bandon, Coos County, Oregon, using hypothetical Cascadia and Alaska earthquake scenarios: Oregon Department of Geology and Mineral Industries Special Paper 43, 57 p.

Funding: This map was funded under award #NA09NW54670014 by the National Oceanic and Atmospheric Administration (NOAA) through the National Tsunami Hazard Mitigation Program.



Nature of the Northwest Information Center 800 NE Oregon Street, #28, Ste. 965 Portland, Oregon 97232 telephone (971) 673-2331 http://www.naturenw.org

Tsunami Inundation Map Index